

## Stability of $\text{Fe}_2\text{O}_3$ in the lower mantle

Dr. Takeshi Arimoto (Postdoctoral Reseacher, GRC)

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### Abstract

$\text{Fe}_2\text{O}_3$  hematite is one of the abundant iron oxide on the Earth. Because of the geophysical importance of  $\text{Fe}_2\text{O}_3$ , phase relations of  $\text{Fe}_2\text{O}_3$  at high pressure have been investigated using a Kawai-type multianvil apparatus (Ito et al., 2009) and a laser heated-diamond anvil cell (e.g., Ono and Ohishi, 2005; Bykova et al., 2016). However, stability of  $\text{Fe}_2\text{O}_3$  in lower mantle condition is not well understood and there are some inconsistencies among their studies. In this study, we made high pressure and high temperature experiments using KMA with sintered diamond anvils in conjunction with in situ X-ray observations up to 60 GPa and 1500 K in BL04B1, SPring-8. Formation of  $\iota\text{-Fe}_2\text{O}_3$  ( $\text{Rh}_2\text{O}_3$  II-type) was observed at ~40 GPa region, which is consistent with data reported by Ito et al. (2009). Furthermore, single phase of  $\zeta\text{-Fe}_2\text{O}_3$  ( $\text{GdFeO}_3$ -perovskite-type) was observed above pressures of ~50 GPa. I will talk about stability of  $\text{Fe}_2\text{O}_3$  and implications for oxygen cycling in the deep part of lower mantle.